



Soil organic carbon density under different agroforestry systems along an elevation gradient in north-western Himalaya

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Abstract

The present study was carried out to understand the variation in soil organic carbon (SOC) density under different agroforestry systems (agri-silviculture, agri-horticulture, agri-horti-silviculture, agri-silvi-horticulture, silvipasture and grasslands) prevailing at three elevations of E₁ (900-1300m; mixed deciduous forest), E₂ (1301-1700 m; chirpine forest) and E₃ (1701-2100 m; oak forest) in Giri catchment of Himachal Pradesh, Himalaya. The results revealed that significantly higher mean SOC was observed in agri-horticulture system (2.22%) as compared to other systems. The bulk density of soil was significantly higher at elevation E₃ (1.28 g cm⁻³) as compared to other elevation ranges. Mean maximum soil organic carbon density upto 30 cm soil depth was observed in grassland (53.45 t C ha⁻¹) followed by agri-horti-silviculture (52.57 t C ha⁻¹), agri-horticulture (51.88 t C ha⁻¹), agri-silvi-horticulture (51.18 t C ha⁻¹) and agri-silviculture (50.01 t C ha⁻¹). Higher SOC density (53.91 t C ha⁻¹) was recorded at elevation E₃.

Keywords: Agroforestry systems, Bulk density, Elevation, Soil carbon density

Introduction

The Clean Development Mechanism (CDM) protocol has highlighted aboveground carbon sequestration through reforestation and afforestation practices, but soils might be serving as largest reservoir of carbon in terrestrial ecosystem (Jha *et al.*, 2003; Hashimoto *et al.*, 2009; Shukla *et al.*, 2017). It has three times higher carbon than vegetation systems of the world and plays a crucial role in carbon cycle (Bohra *et al.*, 2014). Regular addition of leaf litter from aboveground and dead fine roots from belowground has been studied over the past few decades on which soil carbon stock (SOC) depends (Becker *et al.*, 2015; Shukla *et al.*, 2017). It also acts as a carbon sink 1,500-1,550 Pg (Sheikh *et al.*, 2009) and source of carbon depending on the availability and decomposition

changes in organic matter (Dey, 2005). Topographical factors and elevation gradient have an effect on properties of soil through soil erosion and runoff. These processes accelerates the nutrient depletion, decline in soil pH and acidification, build-up of salts in the root zone, nutrient/ elemental imbalance and disruption in elemental cycles; reduction in activity and species diversity of soil fauna, decline in biomass C and depletion of SOC pool (Lal, 2004). Temperature and rain fall varies along with elevation change which impacts soil organic carbon stock through addition of organic matter and changing decomposition rate (Zhang *et al.* 2009).

Carbon management through afforestation and reforestation in degraded natural forests is a useful option, but agroforestry is attractive because the agroforestry systems are promising management practices to increase aboveground and soil C stocks to mitigate greenhouse gas emissions (Makundi and Sathaye, 2004). The C sequestration potential of tropical agroforestry systems in recent studies has been reported to vary between 12 and 228 Mg ha⁻¹, with a median value of 95 Mg ha⁻¹. Therefore, based on the global estimates of the area suitable for the agroforestry (585-1215 × 10⁶ ha), 1.1-2.2 Pg C could be stored in the terrestrial ecosystems over the next 50 years (Albrecht and Kandji, 2003). In India, average sequestration potential in agroforestry has been estimated to be 25 tonnes carbon per ha over 96 million ha (Sathaye and Ravindranath, 1998), but there is a considerable variation in different regions based on biomass production (Dhyani *et al.*, 2009; Ram newaj *et al.*, 2013). However, information on sub-tropical to temperate agroforestry interventions in Giri catchment of Himachal Pradesh on the above discussed parameters of vegetation is meager. Hence, the present investigation was aimed to identify efficient agroforestry systems and to evaluate the agroforestry systems for higher carbon stock in soil pool.

Materials and Methods

Study site and sampling: The present investigation was carried out in sub-tropical to temperate region of Giri catchment during the years 2011 and 2012 under the Department of Silviculture and Agroforestry of Dr. Y.S. Parmar University Nauni, Solan, Himachal Pradesh. Catchment located between 30° 33' 48" and 31° 16' 08" N latitude and 77° 02' 32" to 77° 38' 22" E longitude with an area of 2389 km² (Rao *et al.*, 1989). Catchment is distributed in Shimla, Sirmour and Solan districts of Himachal Pradesh that includes 135 sub-watersheds. Six agroforestry systems were explored for estimation of soil organic carbon densities at each elevation in selected sub-watersheds, which were as follows: S₁: Agri-silviculture system (AS), S₂: Agri-horticulture system (AH), S₃: Agri-horti-silviculture system (AHS), S₄: Agri-silvi-horticulture system (ASH), S₅: Silviculture system (SP), and S₆: Grassland (GS). Stratified random sampling was used for selection of sub-watershed and delineated into three elevations for further investigations *viz.*, elevation E₁ (900-1300 m), E₂ (1301-1700 m) and E₃ (1701-2100 m). Thus, total 39 numbers of experimental sites (treatments) were taken for recording the observations (13 sub-watersheds x 03 elevations).

Soil and data analysis: The composite soil samples from each plot were collected from 0-30 cm depth. Samples were air dried, crushed and passed through 2 mm sieve and stored in cloth bags for analysis. Bulk density of each soil sample was determined by RD bottle method (Singh, 1980). The organic matter in the soil was oxidized by chromic acid (potassium dichromate + conc. H₂SO₄) utilizing the heat of dilution of H₂SO₄. The unreacted dichromate was determined by back titration with ferrous sulphate for organic carbon determination (Walkley and Black method, 1934). The soil organic carbon density was computed by using the formula of Nelson and Sommers, (1996), which was as follows: SOC density (t/ha) = [Soil bulk density (g cm⁻³) x Soil depth (m) x SOC (%)] x 100

The data obtained was subjected to statistical analysis as per the procedure suggested by Gomez and Gomez (1984). Analysis was carried out by using statistical package 'Statistics'.

Results and Discussion

Soil organic carbon: Soil organic carbon (SOC) contents were determined in samples from three elevations selected for each agroforestry system (Table 1). The data revealed that soil organic carbon content was significantly

influenced by the systems. The soil organic carbon (%) in agroforestry systems decreased in order of agrihorticulture (2.22%) > agrihortisilviculture (2.20%) > agrisilvihorticulture (2.17%) > agrisilviculture (2.16%). Significantly higher mean soil organic carbon was recorded in agri-horticulture (S₂: 2.22%) and minimum soil organic carbon was recorded in silvi-pasture system (S₅: 2.08%). However, the SOC obtained in agri-silviculture and agri-silvi-horticulture systems were statistically at par. The difference in soil organic carbon content among different agroforestry systems might be related to difference in structure and composition of vegetation, climatic variation and management inputs.

Table 1. Variation in soil organic carbon (%) relative to elevations and agroforestry systems

Systems (S)	Elevation (E)			Mean (S)
	E ₁ (900-1300 m)	E ₂ (1301-1700 m)	E ₃ (1701-2100 m)	
S ₁	2.17	2.22	2.10	2.16
S ₂	2.22	2.16	2.29	2.22
S ₃	2.21	2.12	2.27	2.20
S ₄	2.12	2.21	2.17	2.17
S ₅	1.98	2.16	2.10	2.08
S ₆	2.26	2.26	2.06	2.19
Mean (E)	2.16	2.19	2.16	
Factors	CD_{0.05}	SEm±		
E	NS	0.020		
S	0.078	0.028		
E×S	0.136	0.049		

Elevation had no significant effect on SOC. However, mean maximum soil organic carbon was observed at elevation E₂ (2.19%) which was higher than E₃ and E₁. The interaction effect of systems and elevations was significant on the soil organic carbon content and revealed that maximum soil organic carbon was observed with S₂E₃ (2.29%) and minimum soil organic carbon (1.98%) was obtained in S₅E₁. Soil is the most effective sequestration reservoir for C in any ecosystems because of the long turnover time of soil organic matter compared with most plant tissues, and because of less inter-annual variability or disturbance-driven losses (Lal, 2004). Generally, soil carbon shares two-third of the total carbon stored by tree based systems. IPCC (2000) has recognized soil organic carbon as one of the five major carbon pools for LULUCF (Land Use, Land Use Change and Forestry). In the present vegetation systems, the SOC varied between 2.08 to 2.22 %. The range of SOC in soils ranging from 1.16 to 3.49 % in grasslands and 1.68 to 3.83 % in chir pine forests were reported earlier

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by Goswami (2014). While in the similar agroforestry systems, Bhardwaj *et al.* (2013) observed SOC contents ranging from 0.89 to 1.22% only.

Soil organic carbon in grasslands is equivalent to tree based systems and is attributed to its intensive root cycling system, which has high content of lignin (Martens, 2000). In the long-term, areas under grasslands have similar potential to store total organic carbon as areas under tree-based systems (Franzluebbers *et al.*, 2000). However, in chir pine silvi-pasture, SOC content was less when compared to other systems. This might be due to two reasons; first in Giri catchment like that of whole Himachal Pradesh, the grasslands and chir pine forests are subjected to frequent fire every year to enhance the forage yield. Hence, the litter in these ecosystems is often burnt. Second, in chir pine silvi-pastures, people collect pine needle litter lying on forest floor every year for its domestic and industrial use. Hence, there is depleted organic matter turnover in these forests, which might have reduced soil organic carbon in chir pine silvi-pasture. Murthy *et al.* (2006) also reported more carbon storage (3.94 t C ha⁻¹) in soils from pure pasture as compared to silvipasture (1.89-3.45 t C ha⁻¹), which was attributed to the high diversity of vegetation.

Table 2. Variation in soil bulk density (g cm⁻³) relative to elevations and agroforestry systems

Systems (S)	Elevation (E)			Mean (S)
	E ₁ (900-1300 m)	E ₂ (1301-1700 m)	E ₃ (1701-2100 m)	
S ₁	1.24	1.10	1.26	1.20
S ₂	1.24	1.12	1.26	1.21
S ₃	1.26	1.13	1.27	1.22
S ₄	1.24	1.12	1.27	1.21
S ₅	1.27	1.15	1.30	1.24
S ₆	1.29	1.19	1.33	1.27
Mean (E)	1.26	1.14	1.28	
Factors	CD_{0.05}	SEm±		
E	0.007	0.003		
S	0.010	0.004		
E×S	NS	0.006		

Bulk density of soil: The mean maximum bulk density of soil was recorded in grassland (1.27 g cm⁻³) which was significantly higher than rest of the systems (Table 2). Among the elevations, the mean maximum bulk density (1.28 g cm⁻³) of soil was observed at elevation E₃ which was significantly higher than elevation E₁ (1.26 g cm⁻³) and E₂ (1.14 g cm⁻³). In the present study, the bulk density of soil in different agroforestry systems decreased in the order: grassland (1.27 g cm⁻³) > silvipas

-ture (1.24 g cm⁻³) > AHS (1.22 g cm⁻³) > ASH (1.21g cm⁻³) > AH (1.21 g cm⁻³) > AS (1.20 g cm⁻³). Similar range of bulk density (0.97 to 1.25 g cm⁻³) in agroforestry systems of Himachal Pradesh was reported earlier (Gowsami *et al.*, 2014; Bhardwaj *et al.*, 2013). They observed that bulk density in grassland varied from 1.07 to 1.22 g cm⁻³ and in silvipasture systems from 0.94 to 1.22 g cm⁻³. But the interaction between systems and elevation exhibited non-significant effect on soil bulk density.

Soil organic carbon density: Elevation and agroforestry systems significantly affected soil organic carbon density (SOCD). Significantly higher SOCD (53.45 t C ha⁻¹) was observed in grassland over silvipasture, agrihortisilviculture and agrisilviculture but at par with agri-horti-silviculture and agri-horticulture (Table 3). Minimum SOCD (49.76 t ha⁻¹) was recorded in silvipasture system. Among elevation, the higher mean SOCD was observed at elevation E₃ (53.91 t C ha⁻¹) which was significantly higher than mean SOCD at elevations E₁ (52.37 t C ha⁻¹) and E₂ (48.15 t C ha⁻¹). The interaction effect of elevation and systems on SOCD revealed that maximum SOCD was recorded in S₃E₃ (56.70 t C ha⁻¹), it was statistically alike with S₆E₁ (56.02 t C ha⁻¹), S₂E₃ (55.64 t C ha⁻¹), S₄E₃ (54.06 t C ha⁻¹) and S₃E₁ (54.04 t C ha⁻¹). Minimum SOCD was recorded in S₂E₂ (47.02 t C ha⁻¹). The SOCD up to 30 cm soil layer under agroforestry systems showed the decreasing precedence as: agrihortisilviculture (52.57 t C ha⁻¹) > agrihorticulture (51.88 t C ha⁻¹) > agrisilvihorticulture (51.18 t C ha⁻¹) > agrisilviculture (50.01 t C ha⁻¹). The effectiveness of agroforestry systems in storing carbon depends on both environmental and socio-economic factors.

Table 3. Variation in soil carbon density (t C ha⁻¹) relative to elevations and agroforestry systems

Systems (S)	Elevation (E)			Mean (S)
	E ₁ (900-1300 m)	E ₂ (1301-1700 m)	E ₃ (1701-2100 m)	
S ₁	51.66	47.20	51.19	50.01
S ₂	52.97	47.02	55.64	51.88
S ₃	54.04	46.96	56.70	52.57
S ₄	51.27	48.20	54.06	51.18
S ₅	48.25	47.92	53.12	49.76
S ₆	56.02	51.57	52.75	53.45
Mean (E)	52.37	48.15	53.91	
Factors	CD_{0.05}	SEm±		
E	1.42	0.51		
S	2.01	0.72		
E×S	3.49	1.25		

In humid tropics, agroforestry systems have the potential to sequester over 70 Mg C/ha in the top 20 cm of the soil and in degraded soils of the sub-humid tropics, agroforestry practices can increase top soil carbon stocks up to 1.6 Mg C ha⁻¹yr⁻¹ (Mutuo et al., 2005). The soil carbon storage in an agroforestry system is regulated by its structural components and their management (Samra and Singh, 2000), soil type (Gupta et al., 2009) and age of the agroforestry plantations (Maikhuri et al., 2000; Saha et al., 2007). Soil carbon density range of 41.05 to 51.04 t C ha⁻¹ in agroforestry systems were recorded by Goswami et al. (2014) and Kanime et al. (2013). In Giri catchment under chir pine forests, Negi and Gupta (2010) observed SOC density (up to 30 cm depth) was varied between 29.71 to 89.03 t C ha⁻¹ with an average value of 57.33 t C ha⁻¹. Likewise, Mahato (2013) and Rajput (2016) also recorded SOC in chir pine silvipasture and grasslands of Himachal Pradesh, which varied between 26.37 to 78.57 t C ha⁻¹ and 31.10 to 43.71 t C ha⁻¹, respectively. Variation in aboveground and belowground biomass carbon (t C ha⁻¹) of vegetation relative to elevations and systems in Giri catchment was also recorded (Table 4 & 5). It was observed that with increase in elevation aboveground biomass carbon increased ($E_1 < E_2 < E_3$), being maximum in E_3 . Similarly belowground biomass carbon increased with increase in elevation ($E_1 < E_2 < E_3$). On the other hand, the maximum aboveground biomass carbon was found in S_5 as compared to others, while the maximum belowground biomass carbon was found in S_4 and the findings were in agreement of earlier report (Chaturvedi et al., 2016).

Table 4. Variation in aboveground biomass carbon (t C ha⁻¹) of vegetation relative to elevations and agroforestry systems

Systems (S)	Elevation (E)			Mean (S)
	E ₁ (900-1300 m)	E ₂ (1301-1700 m)	E ₃ (1701-2100 m)	
S ₁	20.72	23.79	28.44	24.32
S ₂	20.79	23.42	26.06	23.42
S ₃	27.66	29.78	33.98	30.48
S ₄	28.74	29.99	35.96	31.56
S ₅	31.47	35.67	38.81	35.32
S ₆	1.69	1.50	1.51	1.57
Mean (E)	21.85	24.02	27.46	
Factors	CD_{0.05}	SEm±		
E	0.420	0.151		
S	0.594	0.214		
E×S	1.028	0.370		

Table 5. Variation in belowground biomass carbon (t C ha⁻¹) of vegetation relative to elevations and agroforestry systems

Systems (S)	Elevation (E)			Mean (S)
	E ₁ (900-1300 m)	E ₂ (1301-1700 m)	E ₃ (1701-2100 m)	
S ₁	6.33	7.05	8.36	7.25
S ₂	6.31	6.63	7.65	6.86
S ₃	8.30	8.62	10.10	9.01
S ₄	8.73	8.87	10.69	9.43
S ₅	8.70	8.62	10.59	9.30
S ₆	0.93	0.89	0.87	0.90
Mean (E)	6.55	6.78	8.04	
Factors	CD_{0.05}	SEm±		
E	0.106	0.038		
S	0.150	0.054		
E×S	0.260	0.094		

S₁: Agrisilviculture; S₂: Agrihorticulture; S₃: Agrihortisilviculture; S₄: Agrisilvihorticulture; S₅: Silvipasture; S₆: Grassland

Conclusion

Present findings indicated that maximum soil organic carbon density (56.70 t C ha⁻¹) was observed in agri-horti-silviculture system at higher elevation. Therefore, conversion of grassland pasture into agri-horti-silviculture systems is recommended for sequestration of the atmospheric CO₂ in the soil pool at higher elevation. But at lower elevations, grasslands had higher soil organic carbon densities and thus grasslands needs to be conserved and restored.

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